

Conversion of native rangelands into cultivated pasturelands in subtropical ecosystems: Impacts on aggregate-associated carbon and nitrogen

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Abstract: Soil organic carbon (SOC) plays a critical role in the sustainability of grazingland ecosystems around the world. However, maintaining or increasing SOC levels remains a major challenge, particularly in subtropical regions where coarse-textured soils predominate. This study evaluated the long-term (>20 years) impacts of grazingland intensification (conversion of native rangelands into more intensively managed silvopasture and sown pasture) on SOC and nitrogen (N) responses in particle size/density fractions. Treatments consisted of field replicated ($n = 2$) experimental sites that represented a gradient of intensification ranging from native rangelands (low intensification), pine (*Pinus* spp.)–bahiagrass (*Paspalum notatum*) silvopasture (moderate intensification), and bahiagrass pastures (high intensification). Soil organic C and N increased in response to the conversion of native rangelands into more intensively managed grazinglands, but no difference was observed in total SOC and N between silvopasture and sown pasture. Despite the positive impact of intensification on SOC and N pools, accumulation occurred primarily in more labile fractions. For instance, at the 0 to 10 cm depth, light-free C (LF-C) increased from 12.9 g kg⁻¹ soil in the native rangeland to 24.7 g kg⁻¹ soil in the sown pasture. Largest differences between the ecosystems were observed at the 10 to 20 cm depth where LF-C increased by as much as 170% following the conversion from native rangelands to sown pasture. Similar responses were also observed for N. Grazingland intensification showed no effect on soil aggregation, but SOC and N associated with macroaggregates (2,000 to 250 μm) increased with intensification. Results indicate that grazingland intensification promoted SOC and N accumulation, primarily through an increase in the LF fraction.

Key words: aggregates—grazingland—light free—soil organic matter—Spodosol

The impacts of land-use management on soil organic carbon (SOC) and nitrogen (N) dynamics have been documented extensively across different natural and managed ecosystems (Conant et al. 2001; Franzluebbers and Stuedemann 2002, 2013; Shrestha and Stahl 2008; Chan et al. 2011; Schipper et al. 2014). There is a significant body of literature suggesting that adoption of management practices that increase C inputs by increasing plant productivity or reducing C loss through less physical disturbance generally results in a net accumulation of C in soils (Abril and Bucher 1999; Conant et al. 2001, 2003; Chan et al. 2011; McSherry and Ritchie 2013). Although

SOC sequestration rates can be relatively low in grazingland ecosystems, because of the vast area covered and the benefits associated with increasing soil C sequestration, adoption of improved grazingland management practices are of particular importance for the global C cycle. In addition, practices that promote soil C sequestration in grazinglands also tend to result in greater forage production and profitability (Oldeman 1994), and enhance ecosystem sustainability and resilience (Franzluebbers et al. 2014).

Although a significant proportion of the world's grazinglands are currently under pressure to increase productivity (Reid et al. 2004; Follett and Reed 2010), there are still

major gaps in our understanding with respect to the potential impacts of grazingland intensification on SOC and N responses, particularly in the Southern Coastal Plain region of the United States where soils are characterized by sandy texture and often offer limited protection of SOC against mineralization. In addition, a large proportion of grazinglands in this region (particularly native rangelands) is being replaced by more intensive agriculture or urban development. In this context, characterizing changes in SOC and N that occur in response to intensification is a key component of ascertaining the long-term sustainability of grazinglands in this ecoregion.

Research has demonstrated that increased grazingland productivity may not translate into increased SOC stocks, but rather it often affects SOC dynamics in soil size/density pools (Dubeux et al. 2006a). Soil size/density fractions have been extensively used as sensitive indicators of land-use change impacts on SOC and N dynamics (Cambardella and Elliott 1992; Chan 2011; Franzluebbers and Stuedemann 2002; Wen et al. 2013; Badger et al. 2014). Previous research has demonstrated that grazingland intensification often increases C associated with the more labile C pools such as particulate organic C and the light fraction (Franzluebbers and Stuedemann 2002; Conant et al. 2003; Adewopo et al. 2015). Therefore, the separation of SOC pools based on particle size and density is a useful tool to identify the changes that occur in SOC as a function of soil management (Amelung et al. 1998; Baldock and Skjemstad 2000; Romkens et al. 1999; Canellas et al. 2000). This approach has not often been applied to intensively man-

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aged grazinglands, particularly in Spodosols of the Southern Coastal Plain. Soils in this region lack at least two critical mechanisms for sequestering SOC: (1) the physical protection of strongly developed structure and (2) physiochemical protection afforded by soil clay. In addition, the climatic conditions characterized by relatively high temperature and precipitation also limits SOC sequestration in this region. Although our previous studies demonstrated that converting native rangelands into more intensively managed grazingland ecosystems had a positive impact on SOC stocks (Adewopo et al. 2014, 2015; Xu et al. 2015), the effect of intensification on the processes related to SOC accumulation and stabilization in coarse-textured soils has not been fully elucidated. Therefore, the objective of this study was to evaluate long-term changes in SOC and N dynamics in particle size-density pools along a gradient of grazing intensification.

Materials and Methods

Experimental Site. The experimental area was located at the University of Florida, Range Cattle Research and Education Center in south-central Florida (27°23' N, 81°57' W). Mean annual precipitation is ~1,650 mm. Average maximum/minimum temperatures are 28°C/17°C. The soils were classified as Spodosols (Ona and Smyrna fine sands [sandy, siliceous, hyperthermic Typic and Aeric Alaquods, respectively]). The study was conducted on grazingland ecosystems that represent a gradient of management intensities ranging from native rangeland (lowest), silvopasture (intermediate), to sown bahiagrass (*Paspalum notatum*) pasture (highest). The field-replicated sites (~6 ha experimental units) were adjacent to each other but managed independently. Briefly, the management practices used in the native rangeland consisted of periodic burning (every three to four years) and low grazing intensity (~125 animal days ha⁻¹ y⁻¹). Predominant vegetation in the native rangeland consisted of saw palmetto (*Serenoa repens* Bartr.) and a wide variety of grass genera including *Andropogon*, *Panicum*, *Aristida*, and *Schizachyrium* spp. Silvopastures consisted of slash pine trees (*Pinus elliottii*) and Pensacola bahiagrass with moderate management intensification. Nitrogen fertilization history consisted of application of 67 kg N ha⁻¹ y⁻¹ in 1998, 1999, 2001, 2003 to 2007, and 2010 and average stocking rates of 207

animal days ha⁻¹ y⁻¹. Sown pastures represent the typical beef cattle production system in Florida and consisted of rotationally stocked pastures (stocking rate of 360 animal days ha⁻¹ y⁻¹) receiving 67 kg N ha⁻¹ y⁻¹ since 1991 and dolomitic limestone (1 Mg ha⁻¹) in 2001 and 2008. More detailed information about the experimental sites is presented in Adewopo et al. (2014).

Soil Sampling. Five quadrats (20 m × 20 m) spaced ~75 m apart were established along a diagonal transect within each field-replicated experimental unit. In each quadrat, five soil core samples (3 cm diameter × 10 cm height) were collected from randomly selected locations in mid-summer (June and July of 2012) from the 0 to 10 and 10 to 20 cm depths and combined into a composited sample by depth. Soil samples were air-dried and subsequently sieved through a 2 mm screen.

Soil Aggregate Fractionation. Soil aggregate size separation was performed following the wet sieving method described by Six et al. (1998). Briefly, aggregates were separated by wet-sieving air-dried samples through a series of sieves. Aggregate fractions included macroaggregates (2,000 to 250 μm), microaggregates A (250 to 150 μm), microaggregates B (150 to 53 μm), and mineral-associated pool (silt + clay fraction; <53 μm) (figure 1). Each fraction was oven-dried at 55°C and weighed. Although information relative to SOC distribution into the various aggregate size fractions has been reported in our previous work (Silveira et al. 2014), in this current study our main focus was to evaluate the contribution of protected (occluded) versus unprotected (light-free [LF]) pools associated with the different size classes. To accomplish this goal, each aggregate class was further separated into light and heavy fractions following the procedure described by Six et al. (1998) (figure 1). Light-free material, isolated using sodium (Na) polytungstate solution (1.85 g cm⁻³), was transferred to a 20 μm nylon filter, rinsed thoroughly with deionized water, and dried at 55°C. Similarly, the heavy fraction was rinsed with deionized water and then dispersed in 0.5% hexametaphosphate, shaken for 18 hours on a reciprocal shaker, and sieved through either a combination of 250 μm and 53 μm sieves (macroaggregates) or a 53 μm sieve only (microaggregates). Because sand particles are expected to provide little or no protection of SOC, we anticipated that light-free C (LF-C) fraction represented a significant proportion

of C associated with aggregate size fractions. Although our previous studies indicated that intensification promoted C accumulation in the particulate organic matter fraction (Adewopo et al. 2015), it was unclear whether this response was due to accumulation of plant-like material recovered in this fraction versus aggregate-protected C. Therefore, in this current study we decided to further evaluate the impact of nonaggregate (LF) and aggregate-protected (occluded) fractions on C and N accumulation in soils subjected to different management intensities.

Total C and N concentrations in the size/density aggregate fractions were determined by dry combustion using a Flash EA 1112 Series elemental analyzer (Thermo Fisher Scientific Inc., Waltham, Massachusetts).

Statistical Analysis. Data were analyzed using the SAS PROC MIXED procedure (SAS 2001). Soil depth and grazingland ecosystem were considered fixed effects, with replicates considered random effects. The PDIFF test of the LSMEANS procedure and single degree of freedom orthogonal contrasts were used to compare means. Treatments and their interactions were considered significant when F-test *p*-values were <0.05. The means reported are least squares means.

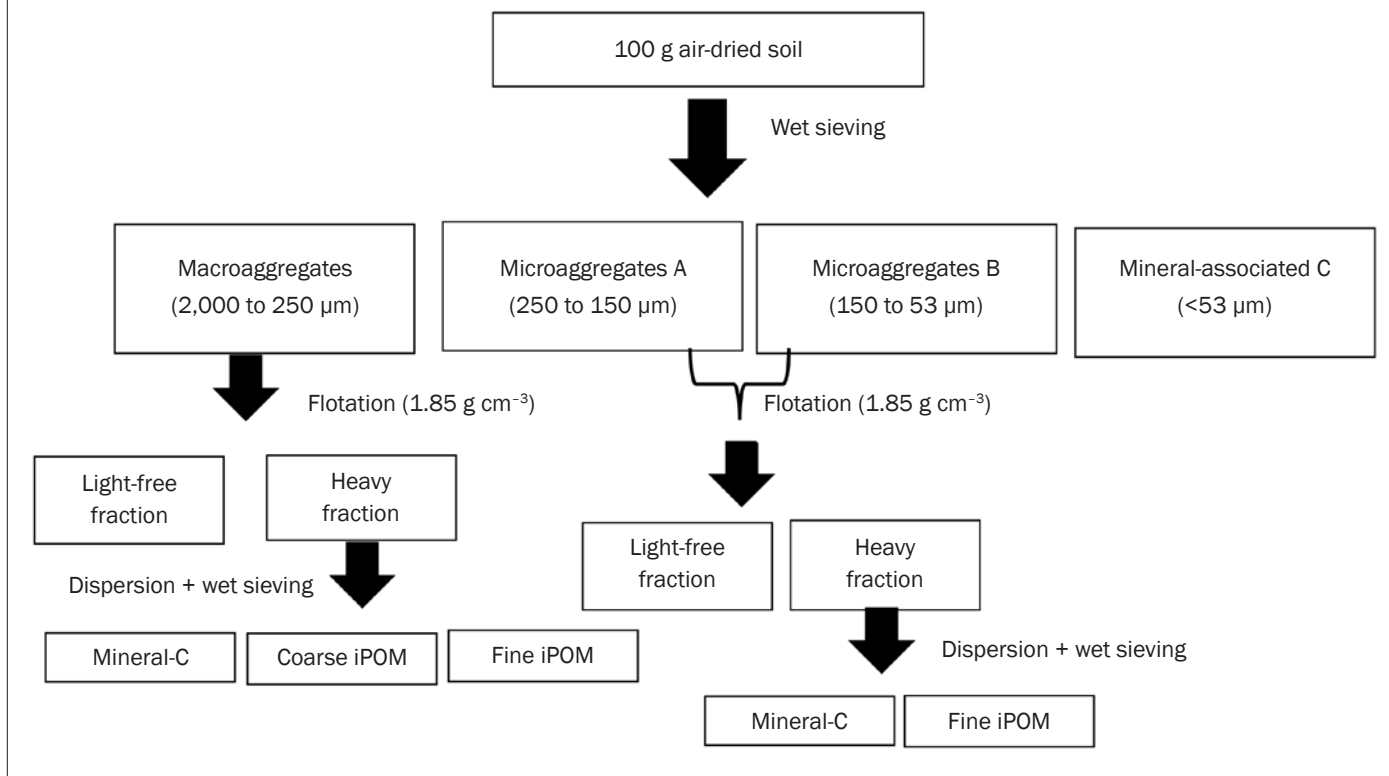
Results and Discussion

Soil Organic Carbon and Nitrogen Distribution among the Various Aggregate Size Classes. As reported in our previous work (Adewopo et al. 2014, 2015; Xu et al. 2015), total SOC and N stocks increased in response to grazingland intensification (table 1). At the 0 to 10 cm depth, total SOC in the native rangeland was ~29% to 37% less than that in the silvopasture and sown pasture, respectively. The pattern of response was observed at the 10 to 20 cm depth. No differences in soil bulk density among the grazingland biomes were observed. Nitrogen concentrations across both depths were, in general, ~57% to 70% smaller in the native rangeland compared with silvopasture and sown pasture system. No differences in total SOC or N concentrations between silvopasture and sown pasture were observed.

Grazingland intensification showed no effect of soil aggregation (figure 2). At the 0 to 10 cm depth, the greatest proportion of soil mass was recovered in the microaggregate A fraction (250 to 150 μm) (average of 382 g kg⁻¹ soil), followed by the microaggregate B (150 to 53 μm) and macroaggregate (2,000

Figure 1

Aggregate fractionation procedure as modified from Six et al. (2000). iPOM is intraparticulate organic matter.



to 250 µm) fractions (average of 348 and 245 g kg⁻¹ soil, respectively). The smallest soil mass was associated with the mineral fraction (<53 µm) (average of 22 g kg⁻¹ soil). At the 10 to 20 cm depth, the greatest soil mass was recovered in microaggregate B size fraction (441 g kg⁻¹ soil), and decreased in the following order: microaggregate B (361 g kg⁻¹ soil) > macroaggregate (176 g kg⁻¹ soil) > mineral (35 g kg⁻¹ soil). Similar results were observed in our previous study using dry sieving separation method (Silveira et al. 2014).

Although no treatment effect was observed on soil aggregation, intensity of grazing land management showed a significant effect on SOC and N distribution among the various aggregate size classes (table 1). In general, major differences in SOC and N among the three land uses were observed in the macro- and microaggregate A pools. At the 0 to 10 cm depth, SOC associated with the macroaggregate fraction was ~36% to 79% greater in silvopasture (12.8 g kg⁻¹ soil) and sown pasture (16.8 g kg⁻¹ soil), respectively, than that in the native rangeland (9.4 g kg⁻¹ soil). Similarly, at the 10 to 20 cm depth, SOC associated with macroaggregates increased from 2.7 g kg⁻¹ soil in the native rangeland to 7.8 g kg⁻¹ soil in the sown pasture.

Nitrogen accumulation in response to grazing land intensification occurred mainly in the macroaggregate pool (table 1). Nitrogen associated with macroaggregates increased two- to four-fold in response to the conversion of native rangelands into sown pastures. No effect of grazing land intensification was observed in the finer (<250 µm) aggregate fractions.

The relative distribution of SOC associated with each aggregate size fraction (percentage of total C in the bulk soil) was not affected by management intensification. On average, at the 0 to 10 cm depth, ~50% of total soil C was present in the macroaggregate fraction, while ~26%, 15%, and 8% of total C was associated with the microaggregate A, microaggregate B, and mineral-associated fractions, respectively.

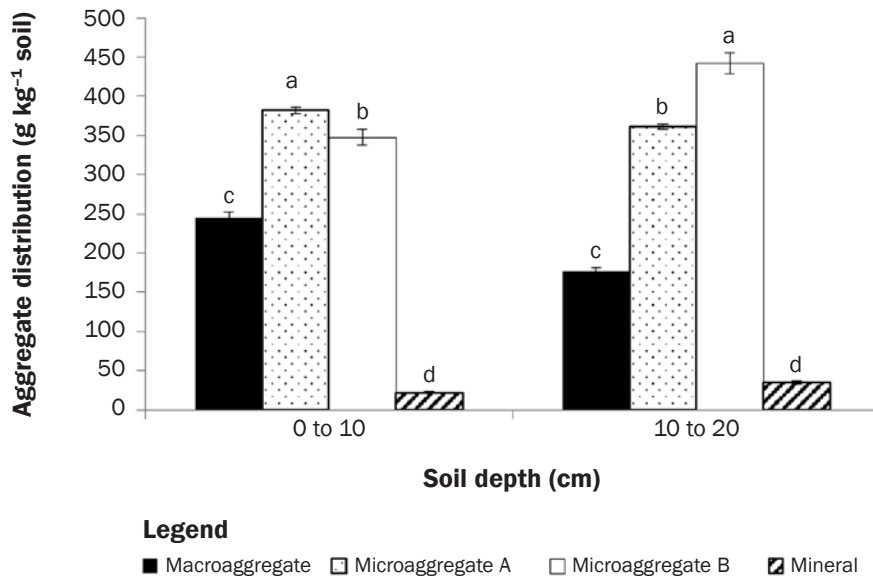
Light-Free Carbon. Grazing land intensification increased SOC and N associated with the LF fraction (table 2). The LF fraction associated with macro- and microaggregates was generally enriched in C and N in the sown pastures compared with silvopasture and native rangelands. Largest differences were observed in the macroaggregate pool. At the 0 to 10 cm depth, LF-C concentrations associated with the macroaggregates were

two- to eight-fold greater in the sown pasture (62 g LF-C kg⁻¹ aggregate) as compared with the silvopasture (27 g LF-C kg⁻¹ aggregate) and native rangeland (7.8 g LF-C kg⁻¹ aggregate), respectively. A similar response was observed for N. Overall, LF-C increased in the following order: native rangeland < silvopasture < sown pasture. Light-free C in both macroaggregate and microaggregate fractions decreased with soil depth.

The mass of C and N associated with LF (g C or N kg⁻¹ soil) followed a similar pattern of response (table 2). Sown pasture exhibited the greatest LF-C (32 g LF-C kg⁻¹ soil) and LF-N (5.9 g LF-N kg⁻¹ soil) compared with the silvopasture (20 g LF-C kg⁻¹ soil and 3 g LF-N kg⁻¹ soil) and native rangeland (15.8 g LF-C kg⁻¹ soil and 1 g LF-N kg⁻¹ soil). At 0 to 10 cm, LF-C represented 82% of total SOC present in the bulk soil, while the proportion of LF-C relative to total SOC was 50% in the silvopasture and 67% in the native rangeland (table 3). Conversely, at the same depth, the proportion of C recovered in the mineral fraction was greatest in silvopasture, least in sown pasture, and intermediate in native rangeland. At the 10 to 20 cm depth, LF-C accounted for 39%, 43%, and 27% of total SOC in the sown pasture,

Figure 2

Aggregate mass distribution. Macroaggregate = 2,000 to 250 μm , microaggregates A = 250 to 150 μm , microaggregates B = 150 to 53 μm , and mineral (silt + clay) fraction = $<53 \mu\text{m}$. Data represent the average across all treatments. Bars represent one standard error. Means with the same letter within soil depth are not different using the LSMEANS procedure ($p > 0.05$).



silvopasture, and native rangeland, respectively. Similarly, a much greater proportion of N was associated with the LF fraction in the sown pasture (85% of total N) compared with native rangeland (60% of total N) and silvopasture (53% of total N) (table 3).

Intraparticulate Organic Matter. Carbon associated with microaggregate intraparticulate organic matter (iPOM) (0 to 10 cm depth) was affected by grazing intensification, but no differences were observed in the macroaggregates' iPOM fractions (figure 3). At the 0 to 10 cm soil depth, the greatest microaggregate iPOM C was associated with sown pasture (1.3 g kg⁻¹ soil), followed by silvopasture (1 g kg⁻¹ soil) and native rangeland (0.7 g kg⁻¹ soil). No treatment effect was observed in the 10 to 20 cm depth. Similarly, the relative proportion of C associated with fine iPOM decreased from 10% of total SOC in the native rangeland to 6% of total SOC in the sown pasture. Because of the relatively low concentrations (below the detection limit of 0.01%), the amount of N associated with the iPOM fractions was considered negligible.

Grazingland Intensification Impacts on Soil Organic Carbon and Nitrogen Dynamics. Grazingland management practices intended to increase forage or animal production such as the introduction of more productive plant species, fertilization, and

grazing intensity and frequency can strongly influence the amounts and characteristics of SOC and N (Conant et al. 2001; Schuman et al. 2002; Wright et al. 2004; Derner and Schuman 2007; Follett and Reed 2010). Results reported in the current study are in agreement with our previous work that demonstrated that conversion of native rangeland into more intensively managed systems can result in increased SOC and N stocks (Adewopo et al. 2014; Xu et al. 2015). These data are also in agreement with Conant et al. (2001) who reported similar SOC responses as result of the adoption of improved grazingland management practices. Soil organic C accumulation in sown pasture occurred primarily due to the replacement of C3 native species with more productive C4 grasses, and adoption of grazing and fertilizer management practices that resulted in greater ecosystem productivity (Adewopo et al. 2014). Similar responses have been reported previously in a variety of grazingland ecosystem types (Fisher et al. 1994; Conant et al. 2001; Wright et al. 2004; McSherry and Ritchie 2013). However, it should be noted that our findings may be most applicable to ecological sites with similar conditions due to the constraints associated with the nondispersal of replicates in space under the studied long-term ecological units.

Light-Free and Mineral-Associated Carbon and Nitrogen. Despite the positive impact of grazingland intensification on SOC and N, our data suggest that SOC and N accumulation following conversion of native rangelands into sown pastures occurred primarily in more labile forms. For instance, our data show that LF-C increased by as much as 2.7 times with increased management intensity. The increases in labile C forms in response to management intensification observed in this current study are in agreement with previous studies in similar soil type and climate conditions (Dubeux et al. 2006a). Those authors found that litter deposition and litter quality increased as grazingland management intensified (Dubeux et al. 2006b, 2006c), and this was associated with an increase in C accumulation in the LF fraction (Dubeux et al. 2006a). Changes in the LF fraction have been highly correlated with residue input (Gregorich and Jazen 1996; Six et al. 1998), and LF has been widely recognized as a sensitive indicator of the effects of land use management (Janzen et al. 1992). Our previous data on litter deposition and composition (Xu et al. 2015) are consistent with the increases in LF-C observed in the sown pasture in this current study. Franzluebbers and Stuedemann (2002) suggested that cattle trampling can also promote plant residue incorporation and, therefore, affect labile SOC pools, particularly at the surface soil depth. Therefore, in addition to the greater SOC inputs via above- and belowground pools, the greater stocking rate associated with the sown pasture may also have favored accumulation of LF-C and N at the 0 to 10 cm depth.

Although no treatment effect on SOC mass (g C kg⁻¹ soil) was observed in the mineral fraction, our results demonstrated that a greater proportion of SOC was associated with the mineral fraction in the silvopasture (0 to 10 cm depth) compared with the other ecosystems. However, no differences in mineral C between silvopasture and sown pastures were observed at the 10 to 20 cm depth. This response contradicts Haile et al. (2010) who observed a greater proportion of SOC associated with mineral fraction in silvopastures versus sown pastures in Florida. These latter authors suggested C associated with the mineral fraction was primarily derived from slash pine trees, and the contribution of the trees increased as the soil depth increased.

Table 1

Soil organic carbon (SOC) and nitrogen (N) distribution among the various aggregate size classes at two soil depths as affected by grazingland management.

| Grazing land ecosystem | Total SOC (g C kg ⁻¹ soil) | Macroaggregate (g C kg ⁻¹ soil) | Microaggregate A (g C kg ⁻¹ soil) | Microaggregate B (g C kg ⁻¹ soil) | Mineral (g C kg ⁻¹ soil) | Total N (g N kg ⁻¹ soil) | Macroaggregate (g N kg ⁻¹ soil) | Microaggregate A (g N kg ⁻¹ soil) | Microaggregate B (g N kg ⁻¹ soil) | Mineral (g N kg ⁻¹ soil) |
|------------------------|---------------------------------------|--|--|--|-------------------------------------|-------------------------------------|--|--|--|-------------------------------------|
| 0 to 10 cm | | | | | | | | | | |
| Native rangeland | 19.2b* | 9.4c | 5.4b | 2.7a | 1.7a | 1.0b | 0.5b | 0.4a | 0.03a | 0.1a |
| Silvopasture | 27.2a | 12.8b | 6.1ab | 5.4a | 2.9a | 1.7a | 0.7a | 0.4a | 0.4a | 0.2a |
| Sown pasture | 30.2a | 16.8a | 8.5a | 3.3a | 1.7a | 2.0a | 1.1a | 0.6a | 0.2a | 0.1a |
| SE | 4.3 | 1.4 | 0.7 | 0.7 | 0.5 | 0.2 | 0.1 | 0.05 | 0.08 | 0.05 |
| 10 to 20 cm | | | | | | | | | | |
| Native rangeland | 10.8b | 2.7b | 2.3b | 2.3b | 3.5a | 0.3b | 0.1b | 0.01a | 0.01a | 0.2a |
| Silvopasture | 15.4a | 4.2a | 2.8b | 5.6a | 2.8a | 0.7a | 0.2a | 0.01a | 0.4a | 0.1a |
| Sown pasture | 20.1a | 7.8a | 6.0a | 4.0ab | 2.3a | 1.0a | 0.4a | 0.4a | 0.1a | 0.1a |
| SE | 3.0 | 1.2 | 0.8 | 0.6 | 0.6 | 0.1 | 0.06 | 0.08 | 0.08 | 0.05 |

Note: SE = standard error.

*Means ($n = 2$) within a soil depth followed by the same letter are not different using the LSMEANS procedure ($p > 0.05$).

Table 2

Grazingland intensification effects on light-free (LF) soil organic carbon (C) and nitrogen (N) pools at two soil depths.

| Grazingland ecosystem | Macroaggregate LF-C (g C kg ⁻¹ aggregate) | Microaggregate LF-C (g C kg ⁻¹ aggregate) | Sum LF-C (g C kg ⁻¹ aggregate) | Macroaggregate LF-N (g N kg ⁻¹ aggregate) | Microaggregate LF-N (g N kg ⁻¹ aggregate) | Sum LF-N (g N kg ⁻¹ aggregate) |
|-----------------------|--|--|---|--|--|---|
| 0 to 10 cm | | | | | | |
| Native rangeland | 7.8c* (7.4)† | 8.1b (5.5) | 16.0c (12.9) | 0.3c (0.3) | 0.4b (0.3) | 0.7c (0.6) |
| Silvopasture | 27.1b (5.1) | 11.1a (8.6) | 38.2b (13.7) | 1.5b (0.3) | 0.7ab (0.6) | 2.2b (0.9) |
| Sown pasture | 62.3a (14.8) | 13.4a (9.9) | 75.7a (24.7) | 4.1a (1.0) | 0.9a (0.7) | 5.0a (1.7) |
| 10 to 20 cm | | | | | | |
| Native rangeland | 7.9b (1.1) | 2.1b (1.8) | 10.0b (2.9) | 0.2b (0.03) | 0.1b (0.09) | 0.3b (0.1) |
| Silvopasture | 10.5a (1.8) | 6.1a (4.8) | 16.6a (6.6) | 0.5a (0.08) | 0.3a (0.3) | 0.8a (0.4) |
| Sown pasture | 11.5a (2.6) | 6.8a (5.2) | 18.2a (7.8) | 0.6a (0.1) | 0.4a (0.4) | 0.9a (0.5) |

*Means within a soil depth followed by the same letter are not different using the LSMEANS procedure ($p > 0.05$).

†Values in parentheses represent the mass of LF-C or LF-N per unit of soil (g LF-C or LF-N kg⁻¹ of soil).

Table 3

Soil organic carbon (C) and nitrogen (N) relative distribution among the various particle size/density pools at two soil depths.

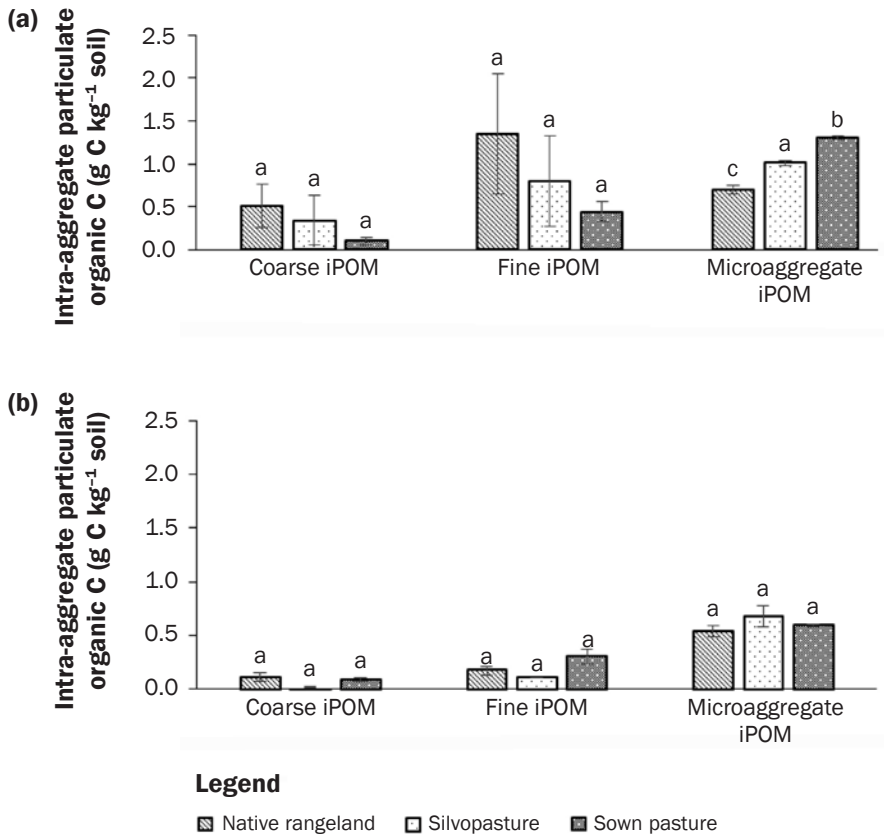
| Grazingland ecosystem | Light-free C (% total C) | Coarse iPOM (% total C) | Fine iPOM (% total C) | Mineral-associated C (% total C) | Light-free N (% total N) | Mineral-associated N (% total N) |
|-----------------------|--------------------------|-------------------------|-----------------------|----------------------------------|--------------------------|----------------------------------|
| 0 to 10 cm | | | | | | |
| Native rangeland | 67b* | 3a | 10a | 20b | 60b | 40a |
| Silvopasture | 50c | 2a | 8a | 40a | 53b | 47a |
| Sown pasture | 82a | 0a | 6b | 12c | 85a | 15b |
| 10 to 20 cm | | | | | | |
| Native rangeland | 27b | 1a | 8a | 64a | 40b | 60a |
| Silvopasture | 43a | 0a | 6a | 51b | 57a | 43b |
| Sown pasture | 39a | 1a | 7a | 53b | 50a | 50b |

Note: iPOM = intraparticulate organic matter.

*Means within a soil depth followed by the same letter are not different using the LSMEANS procedure ($p > 0.05$).

Figure 3

Grazingland intensification effect on intraparticulate organic matter (iPOM) for two soil depths [(a) 0 to 10 cm and (b) 10 to 20 cm]. CiPOM and FiPOM = coarse (2,000 to 250 μm) and fine (250 to 53 μm) intraparticulate organic matter in macroaggregates, respectively. Microaggregate iPOM = fine (250 to 53 μm) intraparticulate organic matter in microaggregates. Bars represent one standard error. Means with the same letter within fraction are not different using the LS-MEANS procedure ($p > 0.05$).



Carbon concentration (g C kg^{-1} aggregate) associated with the mineral pool increased with grazingland intensification. Enrichment of C in this fraction has been reported in various studies under similar environmental and soil type conditions (Dubeux et al. 2006a; Silveira et al. 2013, 2014). Silveira et al. (2013) reported that C concentration in $<53 \mu\text{m}$ fraction can change in response to ecosystem management because the mineral fraction in this soil was dominated by quartz, and, therefore, had low ability to protect SOC against mineralization. Similar to our findings, Adewopo et al. (2015) also reported that conversion of native rangelands into silvopasture systems increased the mineral-associated C (5.7 versus 10 g C kg^{-1} for native rangelands and silvopasture, respectively).

Soil Organic Carbon and Nitrogen Distribution among Various Aggregate Classes. Macroaggregates (2,000 to 250 μm)

accumulated the greatest amounts of SOC and N compared with the other soil aggregate size classes. Averaged across the three ecosystems, macroaggregates accounted for $\sim 50\%$ of total SOC and N at the 0 to 10 cm depth. Similar responses were reported by Sarkhot et al. (2007) who indicated that $\sim 50\%$ of the total SOC was associated with macroaggregate fraction in a Florida Spodosol. These authors concluded that this fraction was more sensitive to management intensification compared with the fine aggregate classes. Corroborating our results, Conant et al. (2004) observed that the majority of the total SOC was associated with the macroaggregate fraction in pasture and forest ecosystems in the southeastern United States. Tonucci et al. (2011) also found that macroaggregates accounted for the greatest proportion of SOC in silvopasture systems. In this current study, SOC associated with

macroaggregates in sown pasture and silvopasture was ~ 1.4 to 1.8 times greater than that in the native rangeland. These results are in agreement with the general assumption that C associated with macroaggregates is more labile (Tiessen and Stewart 1983) than that in the finer particle size pools, and, therefore, more sensitive to ecosystem management (Elliott and Coleman 1988; Six et al. 2004). Similarly, N accumulation due to the conversion of native rangelands into silvopasture or sown pasture occurred primarily in the macroaggregate fraction.

Despite the positive impacts of grazingland intensification on SOC and N associated with macroaggregates, we found that grazingland intensification had no effect on soil aggregation; therefore, SOC and N accumulation observed in the macroaggregate fraction occurred primarily due to increases in the LF pool, particularly at the surface soil depth. Although our previous studies demonstrated that C accumulation in the aggregate size pools was generally promoted as grazingland management increased (Silveira et al. 2014), it was unclear whether this response was due to increases in protected (occluded C) or nonprotected LF. Results from this current study indicated that although C concentration associated with macroaggregates increased from 32 g C kg^{-1} aggregate in the native rangeland to 67 and 70 g C kg^{-1} aggregate in the silvopasture and sown pasture, respectively, LF-C accounted for the majority of the C present in the macroaggregate size class. For instance, at the 0 to 10 cm depth, LF-C represented $\sim 87\%$ of the total C present in the macroaggregate in the sown pasture, while in the native rangeland this proportion was 79%. At the 10 to 20 cm depth, LF-C, on average, accounted for 33% to 41% of total SOC present in the macroaggregates. The hypothesis that SOC accumulation in response to conversion of native rangelands into more intensively managed grazingland ecosystems is primarily due to enrichment in LF pools is further supported by our iPOM data. While our results demonstrated that SOC associated with microaggregate iPOM increased with grazingland intensification, a greater proportion of LF-C was associated with silvopasture and sown pasture compared with native rangelands. Conversely, a much greater proportion of microaggregate C was associated with the mineral pool in native rangeland, suggesting the increases in

SOC associated with increased intensification was caused by the increases in LF-C in sown pasture and silvopasture.

Summary and Conclusions

Long-term grazing land intensification (herein defined as change in vegetation composition, grazing management, and N fertilizer inputs) promoted SOC and N accumulation in subtropical ecosystems. Data also demonstrated that SOC and N accumulation in response to the conversion of native rangelands into silvopasture and sown pasture was primarily attributed to increases in SOC and N associated with the LF pool. Although our results indicated that grazing land intensification affected SOC and N distribution among the various aggregate size classes, unprotected LF comprised a large proportion of SOC and N associated with macro- and microaggregates; therefore, the observed increases in SOC and N may simply reflect accumulation of plant-like material recovered in the various aggregate size classes rather than a mechanism of physical or chemical protection in response to grazing land intensification.

Acknowledgements

This study was partially supported by the Southern Region USDA Program on Sustainable Agriculture Research and Education (SARE).

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